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Rheological behaviour of marine sediments for assessing nautical depth

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Nautical depth concept:

The costs of maintaining the nautical depths in the ports are significant, and authorities issued alternative solutions in order to avoid marine exclusive deposit. These costs will increase exponentially when deeper access channels are to be maintained. In this perspective the nautical bottom concept was developed and implemented in several major ports (Wurpts et al. 2005 - Mehta et al. 2013). This concept is valid because the top layer of the siltation material has, in general, such low strength characteristics that it does not avoid navigation. The expected nautical depth may be guaranteed without removing sediments (Fontein et al. 2006), by using a new concept of dredging system. As suggested for Emden port sediments (Wurpts *et al.*, 2005), the yield shear stress must be lower than 100 Pa for navigability concern.



In the top layer of sediments navigable? siltation material, how making



Site description:

In France, the River Seine carries an average suspended load of about 200 mg/l in the water column, reaching up to 450 mg/l in tidal phases. This suspended material has the nature of quickly entering so-called calm zones as the entrance of Tancarville (France) channel located on the Seine estuary, where the largest retention occurs, leading to intensive dredging.

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Nautical depth and Yield Point from rheometer tests:

The main parameter assumed for navigational purposes is the initial resistance of the mud, when deformed, represented by the yield-strength. Rheological tests were performed in sediment samples collected from Tancarville waterway, involving different solid concentrations, indicated a non-linear increasing of the yield shear stress with the solid concentration from 400g/l to 800g/l ($\Phi = 0,13$ to 0,23) by using two geometries (parallel plate and concentric cylinder). The density involving the suggested yield shear stress of 100 Pa is close to 1.4 t/m³ whereas the threshold value of material flow is close to 1,2 t/m³. So, Tancarville sediment can be considered navigable when its density does not exceed this latter value,



Analysis and matching with results from Marennes d'Oléron bay

To precise the coarse particles effects within mud, the fine/solid fraction ψ is defined: $\Psi = \frac{V_f}{V_g + V_f}$ with Vf : volume of fine particles - (clay-silt) - Vg: grains volume (sand) - Vw: water volume (Pantet et al. 2010) carvile sample



Summary:

Rheometric tests performed in laboratory allowed to assess the suitability of remoulded sediments to reach low shear resistance in order to be navigable. The flow curves show that this mud material will be able to flow when external shear stress exceeds about 30 Pa and the density close to 1,2 t/m3. The dredging method which is applied must shows that the material can be navigable up to a yield point of 100 Pa if its density is lowered to less than 1.4 t/m3. Previous rheological measurement sensitivity allowed to identify five typical sediments that correlate with solid fraction and fine fraction. From the provided diagram, the Tancarville sediment falls in the area of a fluid with threshold,

To understand the relationship between the mud sedimented layer ang the mud suspension layer, further experiments involving a chanel with cohesive sediments are under design to complete the rheological characterization.

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Rheological measurements to improve the understanding of suspended sediment dynamics in the Ems estuary

MUDNET Conference – Delft 29.3. – 30.3.2021

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The Ems estuary



the bed of the lower Ems estuary is temporarily covered with fluid mud layers up to serveral meters high variability of hydrodynamic conditions on different timescales

- lead to complex behaviour of fluid mud concerning formation, consolidation, entrainment and advective transport processes
- goal: to conduct further research to examine the rheological properties of estuarine muds under different hydrological conditions

Density and yield stress measurements in the estuary



- bfg Bundesanstalt für Gewässerkunde
- Rheotune measurements at Emskm 7,2 (upper estuary) on 17.6. and 16.7.2020
- Multiple profiles with Stema Rheotune probe during ebb tide represent different hydrodynamic conditions
- Denisty (black line) and yield stress(red line) profiles differ significantly
- Question: how do sediment/mud samples differ concerning their (rheological) properties, especially thixotropic behaviour?





0	1100	1200	1300
	De	nsity [kg/m³]	

100

Rheological investigation of mud samples







First tests carried out in 2020 with Haake Mars Rheometer

Further investigation of van Veen grabber samples, Emskm 7,2 on 16.7.2020

Systematic investigation of fluid mud samples taken under different hydrological conditions will take place in 2021



IDRM 3 U-tube system

The IDRM is an online 24/7 selfpriming pipe rheometer with an arrangement of U-tubes. It enables determination of the rheogram based on simultaneous measurements of pressure differentials and flowrates in vertically oriented tubes with a variety of diameters. Main output of the system is a rheogram, including Bingham parameters and density every 10 minutes!

<u>IHC Deltares Online Pipe Rheometer (IDRM)</u>

Water management thickener by rheological control of the underflow



BHP and EXPANDE Mining Open Innovation Program of Fundación Chile

Rabinovich-Mooney Transformation





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Rheological analysis of a model mud for laboratory erosion experiment Pierre Lecostey¹, Guillaume Gomit¹, Sébastien Jarny¹, Lionel Thomas¹ ¹ Institut P', CNRS, Université de Poitiers, ENSMA, Boulevard Marie et Pierre Curie, Futuroscope 86962, France

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Context and Purpose

Rheological properties of mud have a crucial impact on their erosion. To better understand this link, a local study of the erosion process is proposed. To make it possible, a clear model mud has been created with a mix of laponite and CarboxyMethylCellulose (CMC) (figure 1). Indeed erosion experiments, using optical technics like Particle Image Velocimetry (PIV) and sediment surface estimation, required clear mud. Furthermore, model mud takes advantage to be rheologicaly stable and reproducible.



Rheological analysis



Flowing rheological tests have been conducted on the model mud and its dilutions. It appears that the sediment can be modelized by an Heschel-Bulkley fluid with a yield stress r_s (figure 2). Herschel-Bulkley model is often used to describe natural fine-grain sediment [1]. The clear mud has a yield stress around 10 Pa which is consistent with real fluidized mud. The dilution of the model mud will be used to create several $\hat{\mathbb{E}}$ laboratory sediments with different Herschel-Bulkley characteristics. Yield stress of each dilution have been represented on the figure 3. The growth of τ_s can be modelized by a power law which is in agreement with natural mud [2].



MUD

Erosion experimentation

The study of the model mud diluted under turbulent flow has been performed in a squared In study of the model mud diluted under turbulent flow has been performed in a squared closed channel (figure 4). Particle Image Velocimetry (PIV) coupled with a surface detection algorithm have been used. The use of the PIV allows the estimation of the bottom shear stress of the flow, important factor of sediment erosion [3], and the mud surface. In addition, the PIV could give more information on the flow like hairpin structure which is known to affect non-fine sediment erosion [4]. The correlation of velocity field data from the PIV with the surface detection allows the analysis of the relationship between the flow structures and the bed sediment motion.



The figure 5 shows a example of the mud surface deformed by the flow just before erosion process. The surface detection code result is plotted in red and the white vectors represent the velocity calculated with the PIV technics.

Fig. 5 PIV vector coupled with surface detection on the model mud



Prospects

- · Link rheological constants with flow characteristics
- · Analyse surface oscillations preceding the erosion of the mud
- · Correlate surface movement to flow structures
- · Study model mud in an open channel

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4 Hydro-sedimentary channe.





1.

2.

3.









A rheological model for mud which avoids the use of a yield stress

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Abstract

A Cross model has been applied to mud samples from the Ems Estuary with different sediment volume fractions ϕ_s . It yields excellent approximations in the area where shear is low. This model represents a transition from a high to a low viscosity and avoids the use of a yield stress. The model parameters strongly depend on ϕ_s . For CFD simulations, the dependency on ϕ_s can be included into the model. The model can then be applied for water, fluid mud and immobile mud.

Introduction and Motivation

Mud samples show a shear-thinning, thixotropic behavior with two-step yielding [1]. Figure 1 shows the scheme of a flow curve and viscosity curve for a mud sample. A Bingham or any other model that consists of a yield stress, e.g. $\tau=$ $\tau_{y,B} + \mu_{\infty}\dot{\gamma}$ results in an infinitely large viscosity for $\dot{\gamma}=0$. Modern rheometers detect small movements below the yield stress $\tau_{\gamma, s},$ which indicates a high but not infinite viscosity.

Thus, the approximation of mud data with a Bingham model is weak at areas of low shear and leads to a possible division by zero in numerical simulations. In natural channels, the velocity shear is typically low. especially when the sediment concentration is high. An improvement is possible with a model which describes a transition from a high to a low viscosity instead. Suitable models are the Cross model or Carreau-type models.



Stress sweep tests were performed with four mud samples ($\varphi_{\scriptscriptstyle S}=0.16~0.13~0.10$ and 0.07, see Figure 2) using the stress-controlled mode of the rheometer.

The Cross model follows the equation

 $\mu = \frac{\mu_0 - \mu_\infty}{1 + (t_r \dot{\gamma})^n} + \mu_\infty$

(1) where $\mu \rightarrow \mu_0$ for $\dot{\gamma} \rightarrow 0$ and $\mu \rightarrow \mu_{\infty}$ for $\dot{\gamma} \rightarrow \infty$ [2]. We wish to apply the model up to an intermediate viscosity μ_1 instead of μ_{∞} :

The results of Equation (2) with the parameters

for the four mud samples are shown in Table 1 and Figure 3. The parameters $\mu_0\,,\,\mu_1\,$ and \mathbf{t}_r increase with increasing $\boldsymbol{\varphi}_s.$ As the parameter nwas close to 1 for all samples, the value was set

to 1. The model is in excellent agreement with

measurement data. A dependency for the

parameters $\mu_0(\varphi_s),\,\mu_1(\varphi_s)$ and $t_r(\varphi_s)$ can be found. The model is valid for $0 \leq \dot{\gamma} \leq \dot{\gamma}_{c}$

Materials and Methods

 $\mu = \frac{\mu_0 - \mu_1}{1 + (t_r \dot{\mathbf{y}})^n} + \mu_1$

(2) The parameter t_r has the unit seconds. The value $\frac{1}{t_r}$ is the value where the transition from μ_0 to μ_1 starts. The parameter n indicates how quick the transition

 $\begin{array}{ll} = 0.16 & \phi_s = 0.13 & \phi_s = 0.10 & \phi_s = 0.07 \\ 1264 \, {\rm kg} \, {\rm m}^3 & \rho = 1215 \, {\rm kg} \, {\rm m}^3 & \rho = 1165 \, {\rm kg} \, {\rm m}^3 & \rho = 1116 \, {\rm kg} \, {\rm m}^3 \end{array}$

Figure 2. Mud samples from the Erns Estuary (Jerngum). Rheological experiments were performed using a Physica MCR301 rheometer. A parallel plate geometry with a gap of 1 mm was used.

4.

5.

Results and Discussion

occurs.

Equation (2).

Table 1. Results of the parameters for the Cross model for different mud samples. "There is an interesting relation between μ_0 and t_{Γ} . The value of the yield stress $\tau_{y,s}$ can be analytically derived as $\tau_{y,s} \approx \frac{\mu_0}{\tau_c}$."

For evaluation, the data between $0 \le \dot{\gamma} \le \dot{\gamma}_c$ was

used. The measurement data was then fitted to

Sample	μ	μ	tr	$\tau_{y,s} \approx \frac{\mu_0}{t_r}$	R^2
	[Pa s]	[Pa s]	[S]	[Pa]	
J 0.16	$3.0\cdot 10^5$	5	2420	124	0.988
J 0.13	$9.0\cdot 10^4$	2	1765	51	0.9798
J 0.10	$2.5\cdot 10^4$	1	1470	17	0.9964
J 0.07	$6.0 \cdot 10^2$	0.3	600	1	0.997



Figure 3. Measurement data and Cross model for the mud samples. The viscosity curve (right) is calculated with Equation (2). The flow curve (left) is then calculated with $\tau = \mu \dot{\gamma}$.

Conclusions

References

This study shows a simple rheological model for mud which avoids the use of a yield stress and is valid in areas of low shear rate. The model was successfully applied to mud samples with different sediment volume fractions. In CFD simulations, such a yield stress-free model has the advantage that it avoids the possible division by zero. Furthermore, it includes the rheological measurement data at very low shear rates which is usually neglected.

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Rheology of cohesive suspensions: A preliminary approach towards the sediment transport

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Introduction

The sediment transport is of major importance as it contributes in the morphology of riverbeds and it may be associated to risks (e.g., floods, scouring). One of the main challenges is to relate the sediment transport rate to the hydraulic conditions applied to the granular bed. However, a fine proportion of cohesive materials may strongly affect the sediment transport by increasing the particle threshold of motion [2, 4]. The present work aims to combine both the rheological measurements and the laboratory experiments on the sediment transport in a flume with clay-sand mixtures to highlight the relationship between the rheology and the transport of such materials. Here, we present the preliminary results on the rheology of pure illite and kaolinite suspensions over a wide range of volume fractions, using both a cone-plane and a novel vane geometry.

Setup and validation



Figure 1: (a) Vane geometry and (b) measurement protocols.

The **rheological measurements** are performed using both a Physica MCR 501 rheometer and a RheolabQC rheometer from Anton Paar, equipped with a cone-plane and a novel vane geometry [Fig. 1(a)], respectively. Three different protocols are used to ensure the reliability of measurements [Fig. 1(b)]. **Cohesive suspensions** are made of illite or kaolinite clay suspended in tap water, and they are prepared at least one day before the rheological measurements. The masses of water m_w and of clay m_c are weighted accurately to estimate the volume fraction of suspensions as

$$\phi = \frac{\rho_w m_c}{\rho_w m_c + \rho_c m_w},\tag{1}$$

where $\rho_w = 1000 \text{kg.m}^{-3}$ is the water density, and $\rho_c = 2600$ and 2800kg.m^{-3} are the densities of kaolinite and illite, respectively. The volume fraction is varied in the range $\phi = [0: 0.269]$ with the cone-plane, and $\phi = [0.147: 0.399]$ with the vane geometry.

Role of the volume fraction

As mentioned above, the rheological behavior of cohesive suspensions is mainly controlled by the volume fraction ϕ (Fig. 3), as already reported by previous studies [1, 2]. The increase of ϕ promotes larger shear stresses τ and apparent viscosities η , for a given shear rate $\dot{\gamma}$. Additionally, the flow curve of suspensions is well described using the Herschel-Bulkley model (dashed lines). The evolution of rheological parameters with ϕ can therefore be addressed, and in particular, the yield stress τ_y quantifying the mobility of mixtures for the purpose of investigating the sediment transport.



Figure 3: (a) Shear stress τ and (b) apparent viscosity η as a function of the shear rate $\dot{\gamma}$ for kaolinite suspensions with different ϕ . The dashed lines are the best fits of the Herschel-Bulkley model.









Figure 4: Yield stress τ_y as a function of ϕ for illite (green) and kaolinite (gray) suspensions, measured by the cone-plane (triangles) and the vane (circles) geometry with different protocols (light to dark colors). Opened symbols are data from the literature. The dashed line is $\tau_y = a \exp(b\phi)$, with a = 1 and b = 17.

As expected, τ_y increases over several decades with increasing ϕ (Fig. 4). At large ϕ , the experimental data may be described by the empirical law $\tau_y = a \exp(b\phi)$ [1], with a = 1 and b = 17 (dashed line), regardless of illite and kaolinite suspensions. At low ϕ , however, the yield stress is overestimated by the empirical model. Finally, our experimental data are included in the scatter of results from the literature (see opened symbols).

Conclusion

- Rheological measurements on pure illite and kaolinite suspensions with various volume fractions are performed using both a cone-plane and a novel vane geometry.
- The rheological behavior of cohesive suspensions is mainly controlled by the volume fraction, while the geometry used causes a bias of measurements.
- The yield stress increases with ϕ , in agreement with empirical laws and data from the literature.
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